

Constraints on Fracture Distribution in Geothermal Fields Using Seismic Noise Beamforming

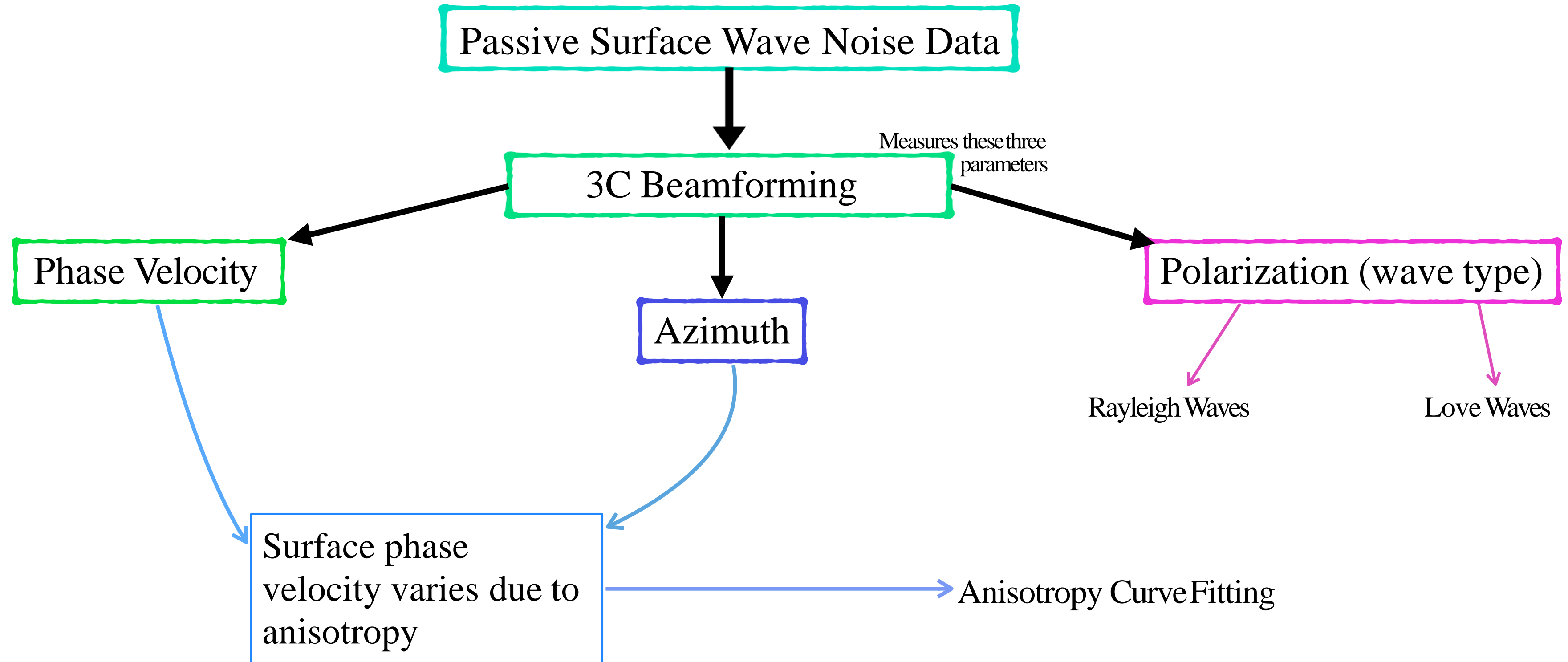
SPE Seismic 2022

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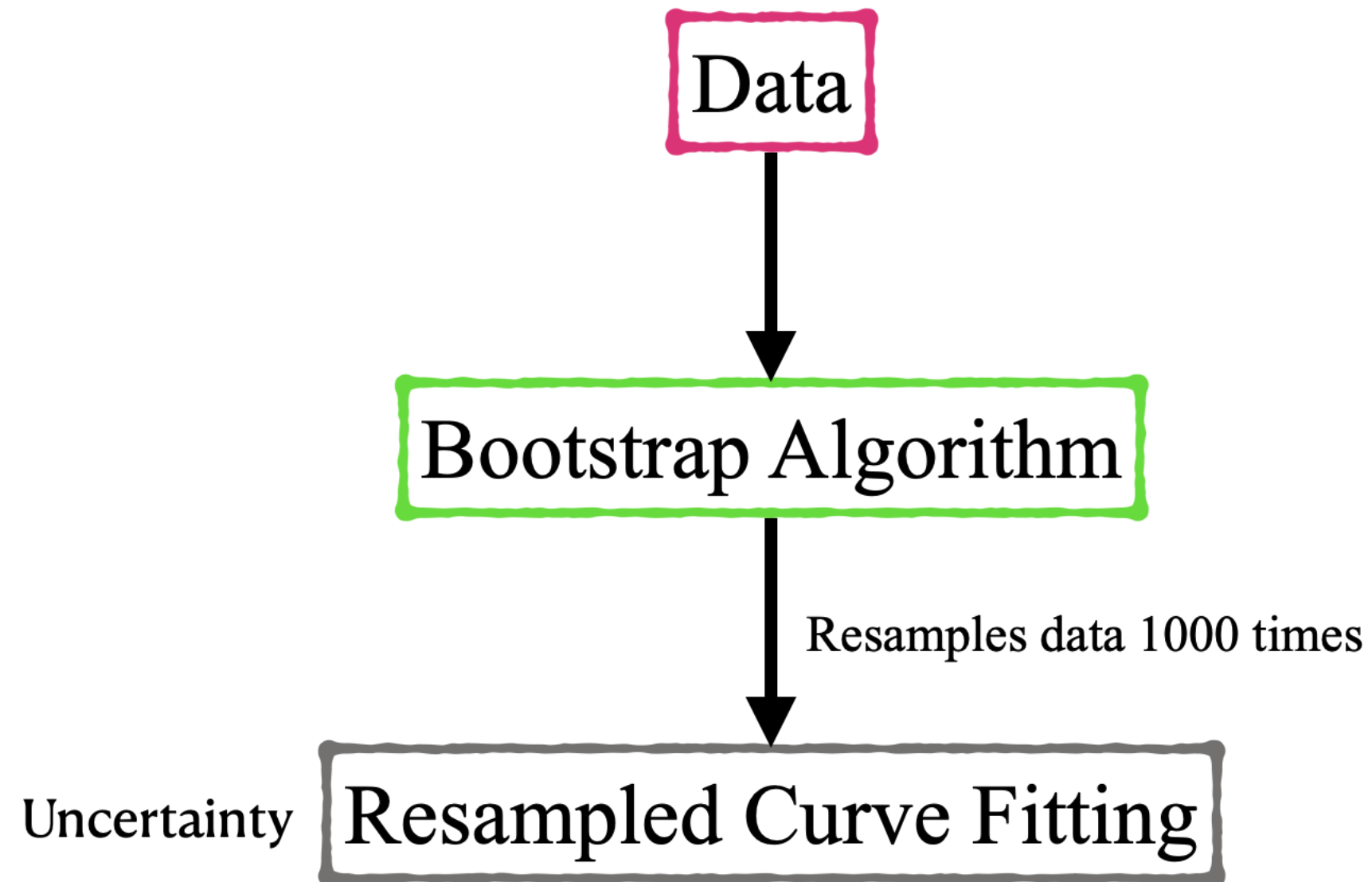
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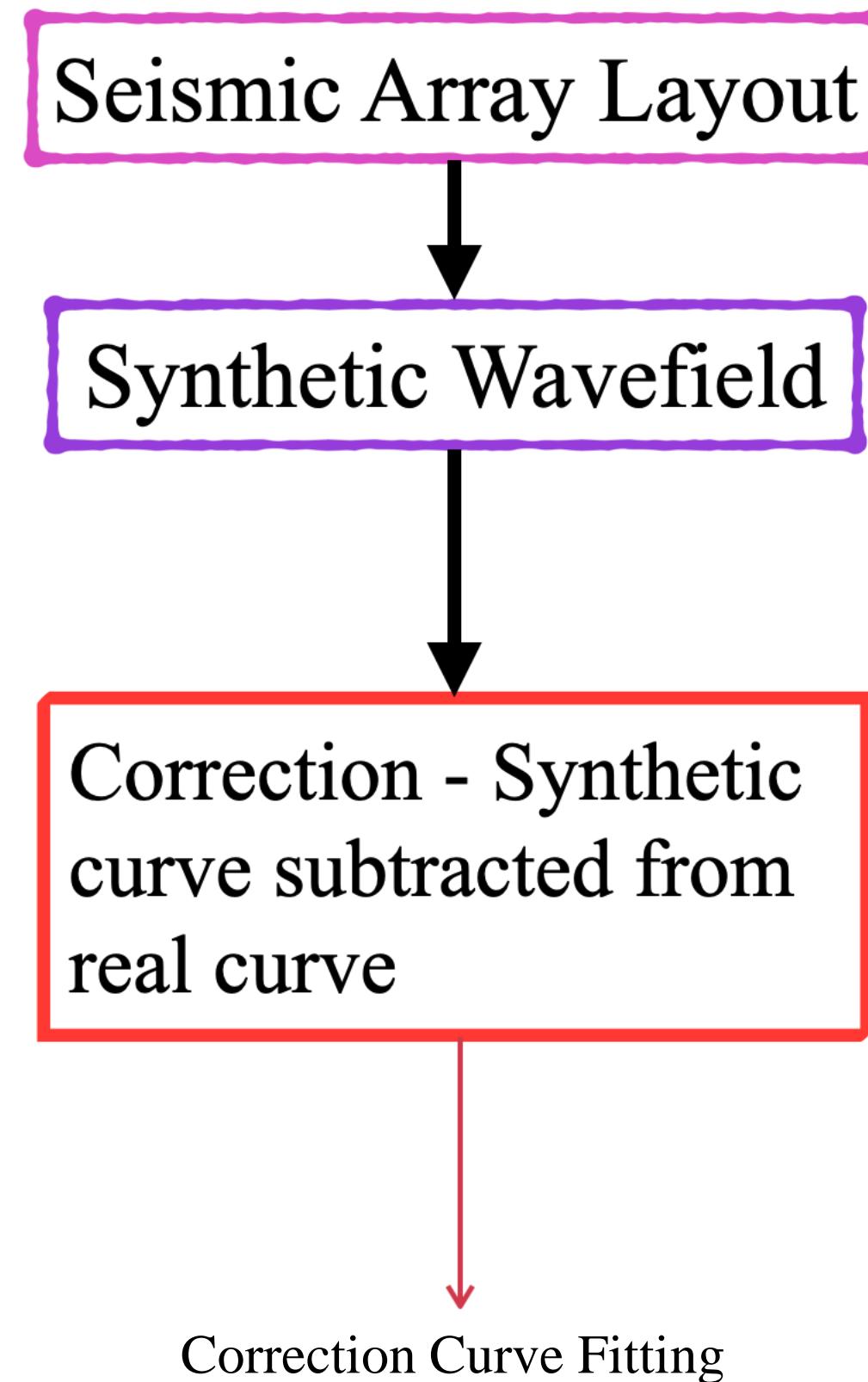
3C Beamforming



Uncertainty Assessment



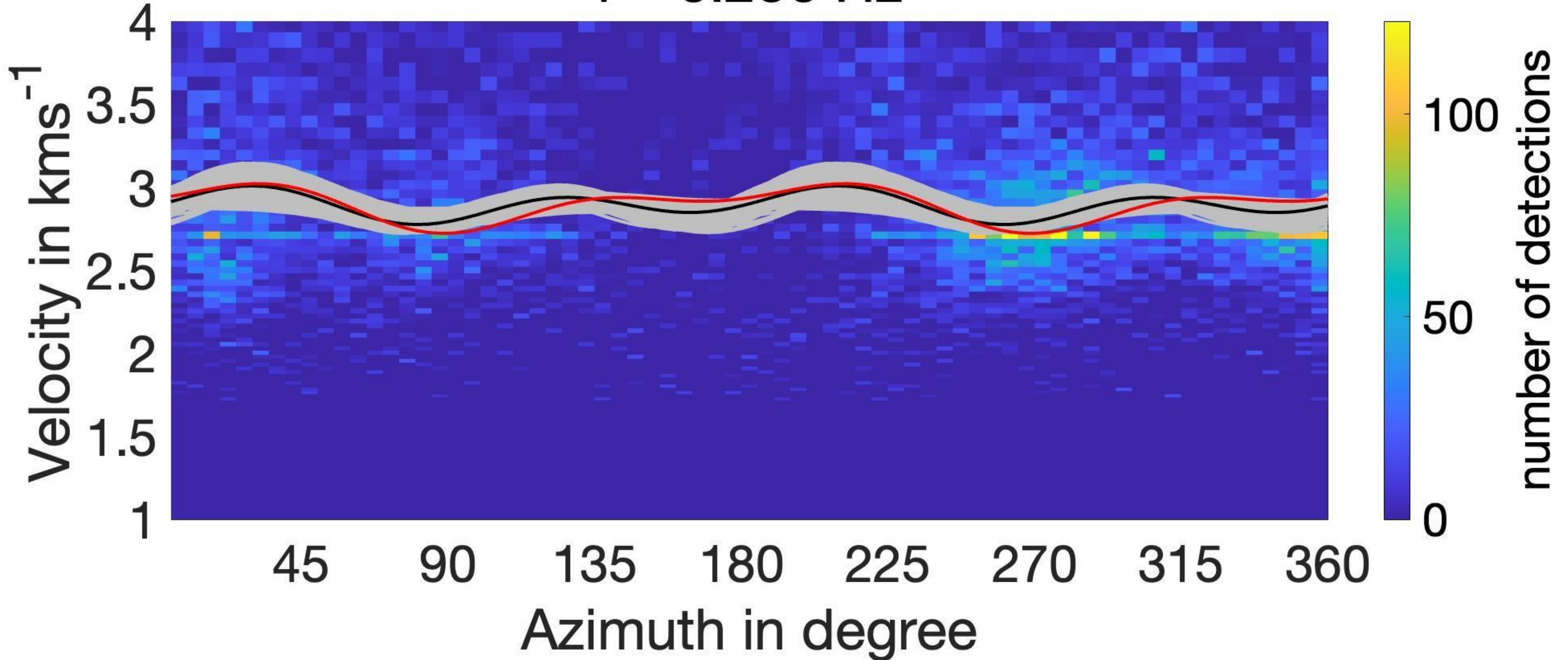
Correcting for Array Effect



Anisotropy

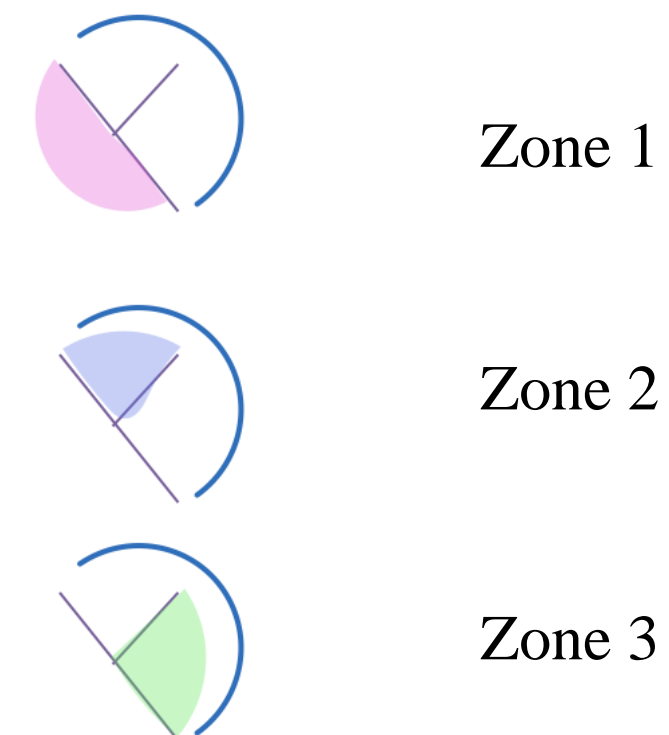
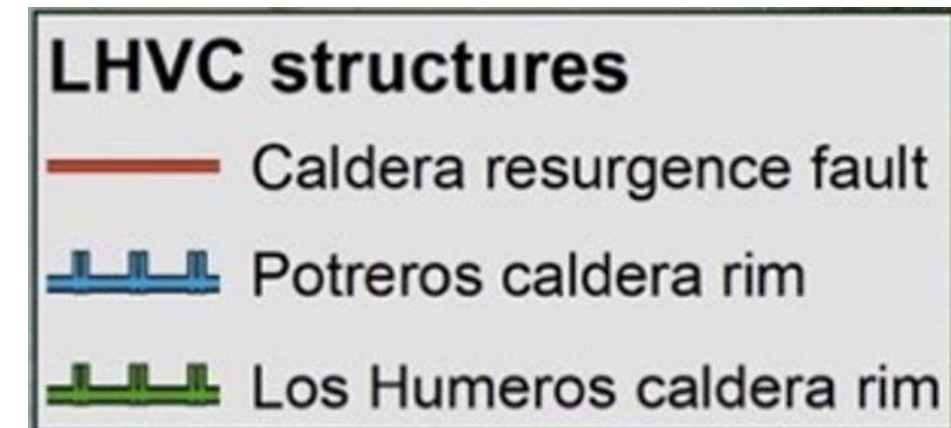
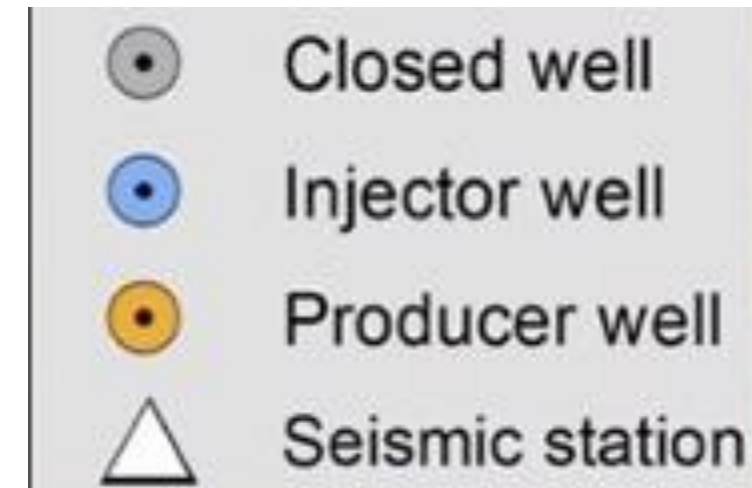
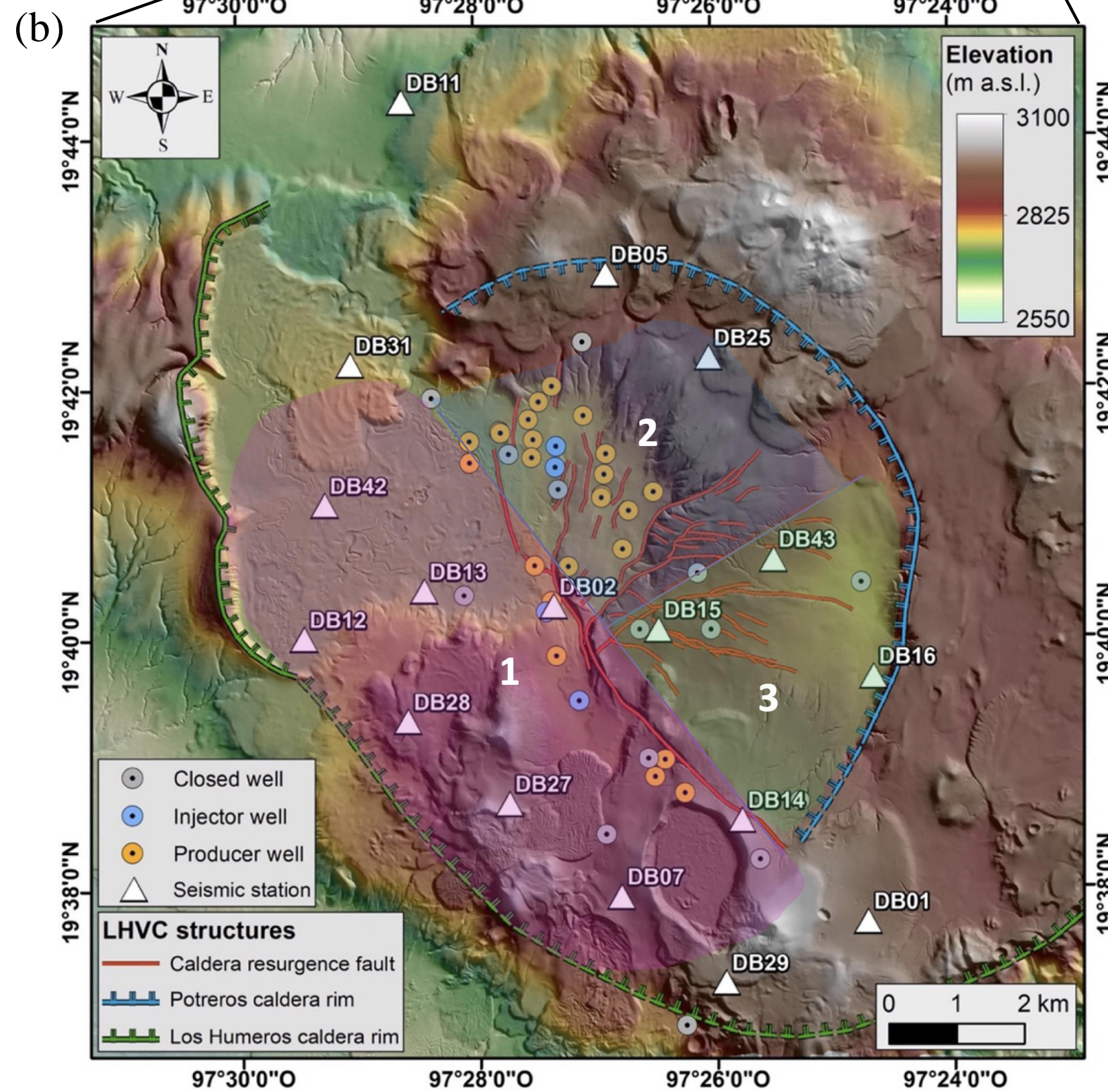
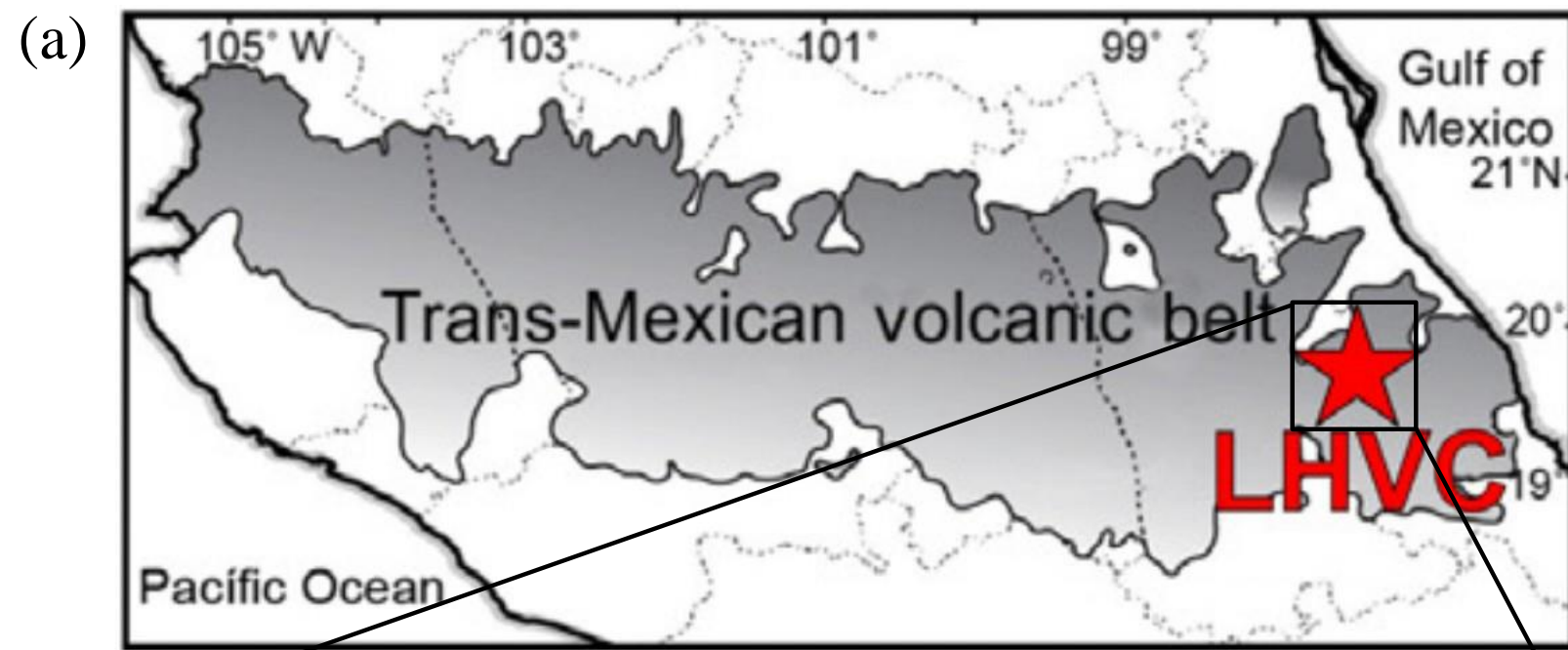
Retrograde Rayleigh Wave

$f = 0.250$ Hz



— Corrected Anisotropy — Mean Anisotropy — Uncertainty

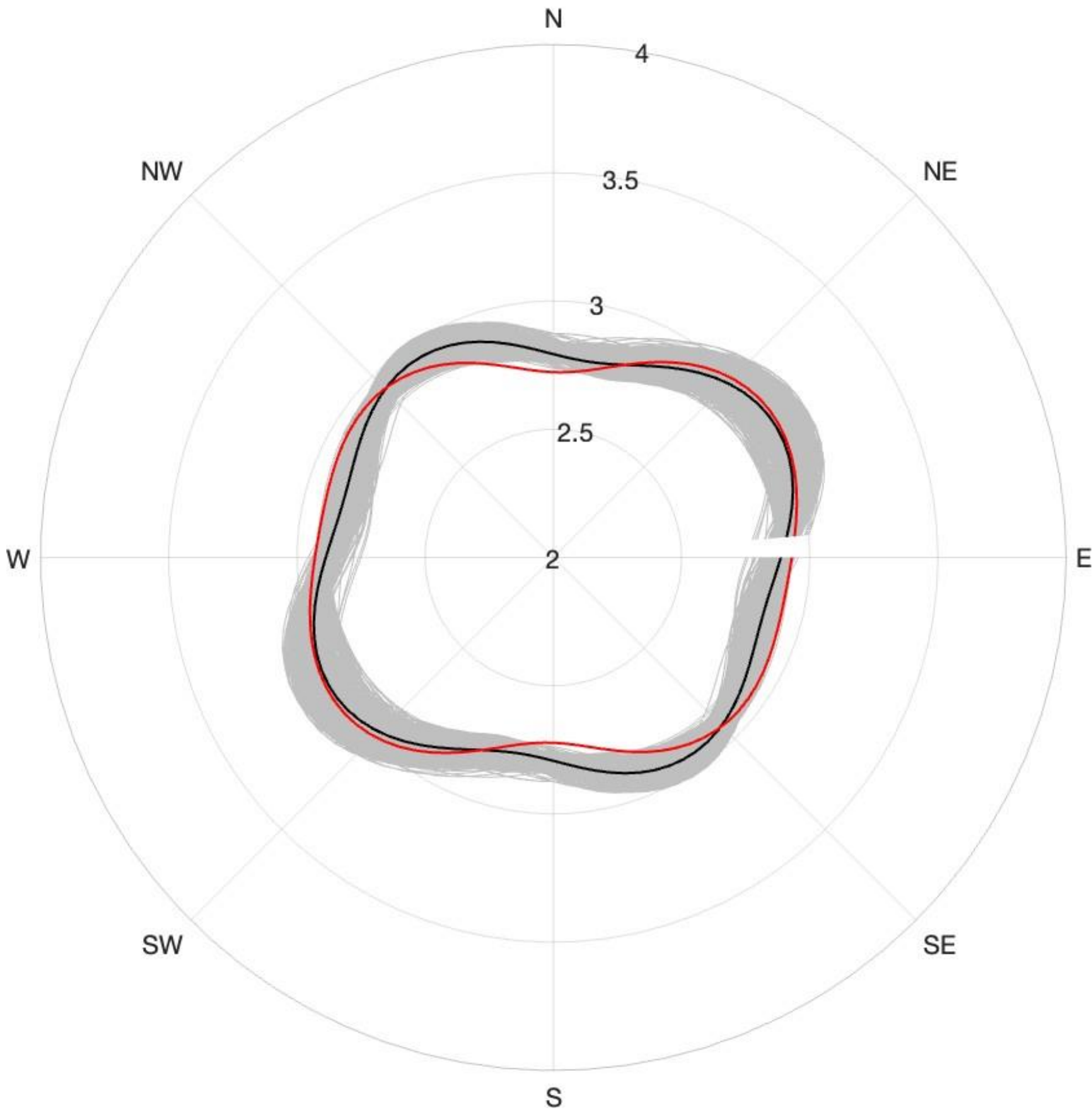
Los Humeros Geothermal Field



(a)

Retrograde Rayleigh Waves

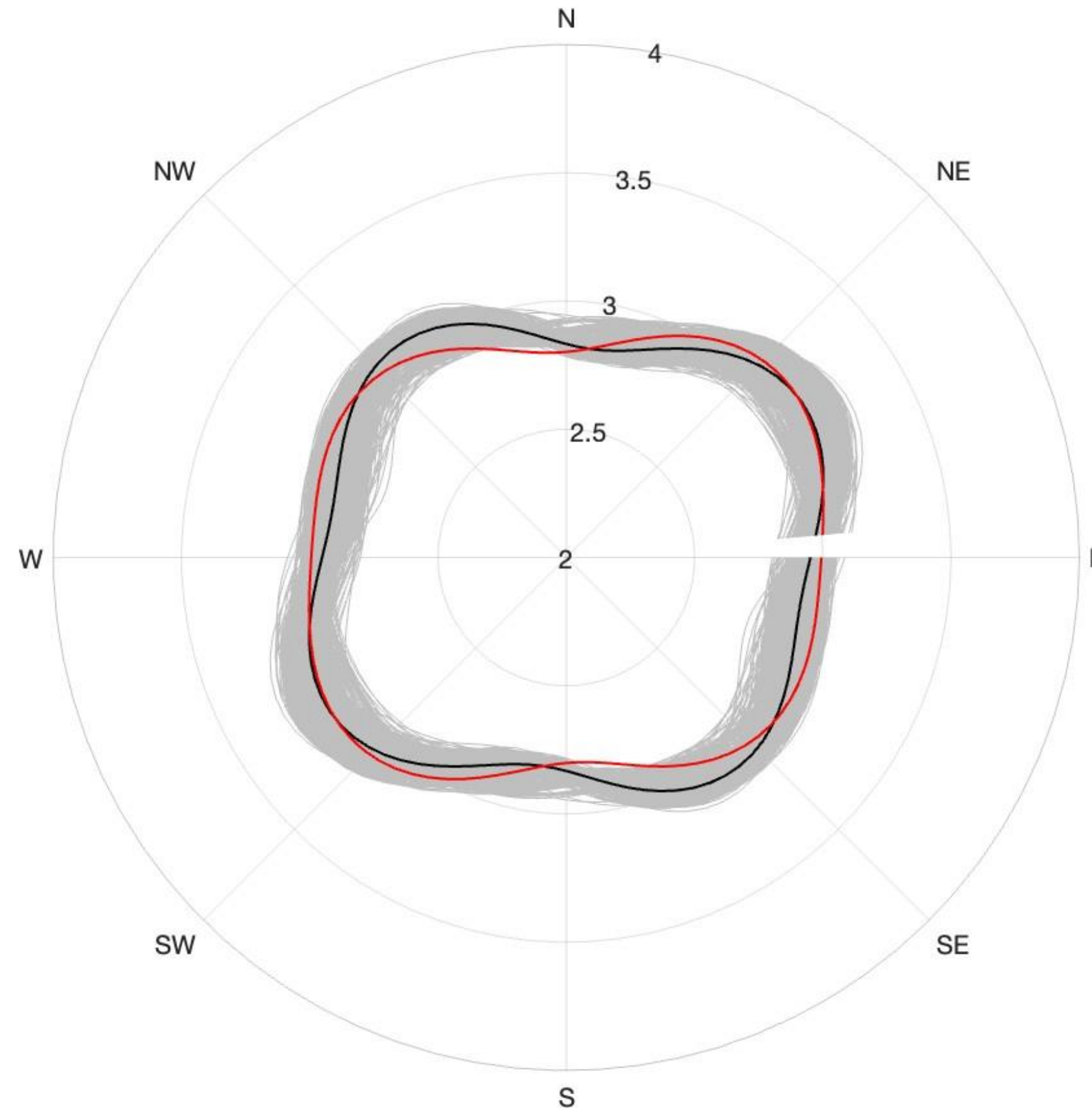
f = 0.250 Hz



(b)

Love Waves

f = 0.250 Hz



Corrected Anisotropy

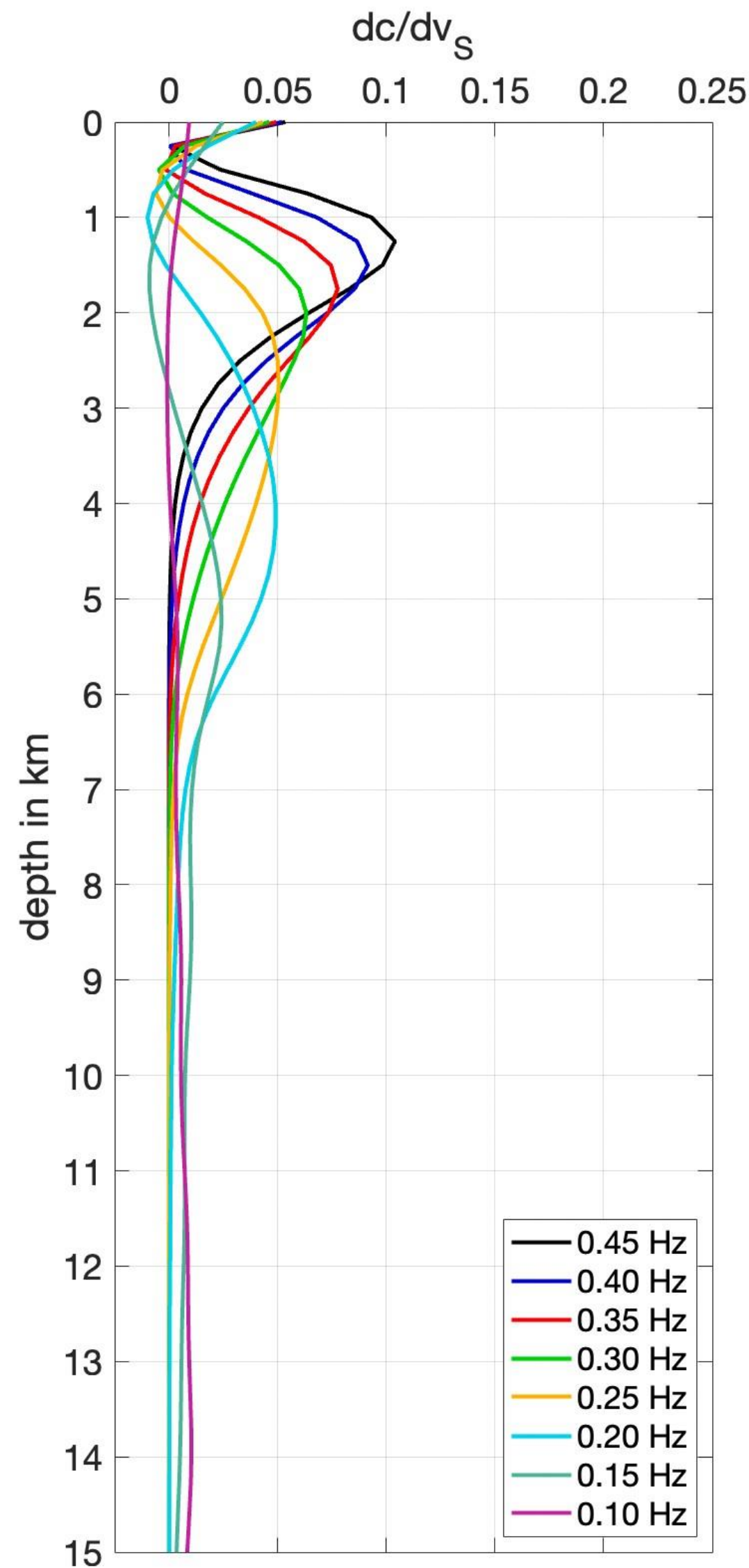


Mean Anisotropy

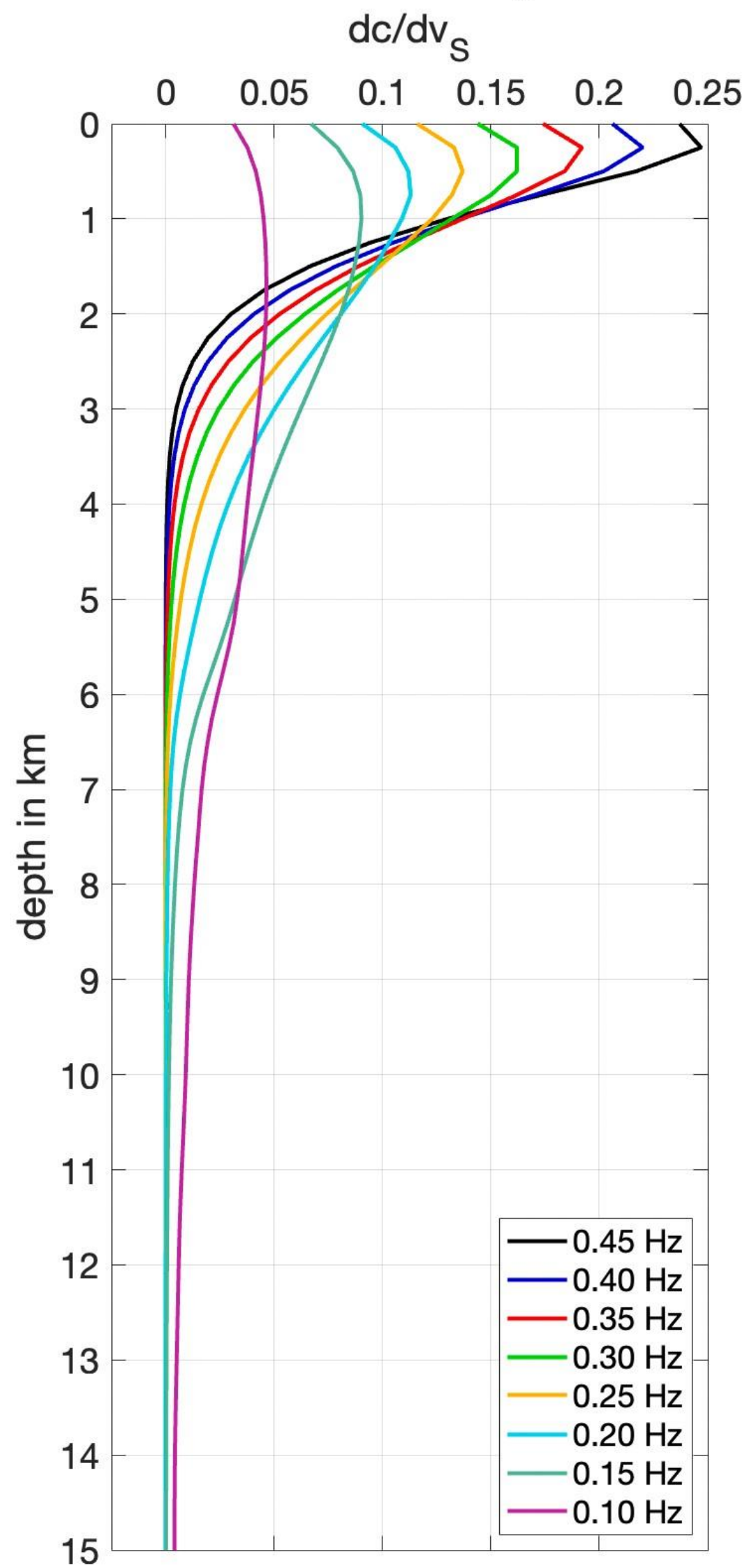


Uncertainty

(a) **Rayleigh Wave Sensitivity Kernels**



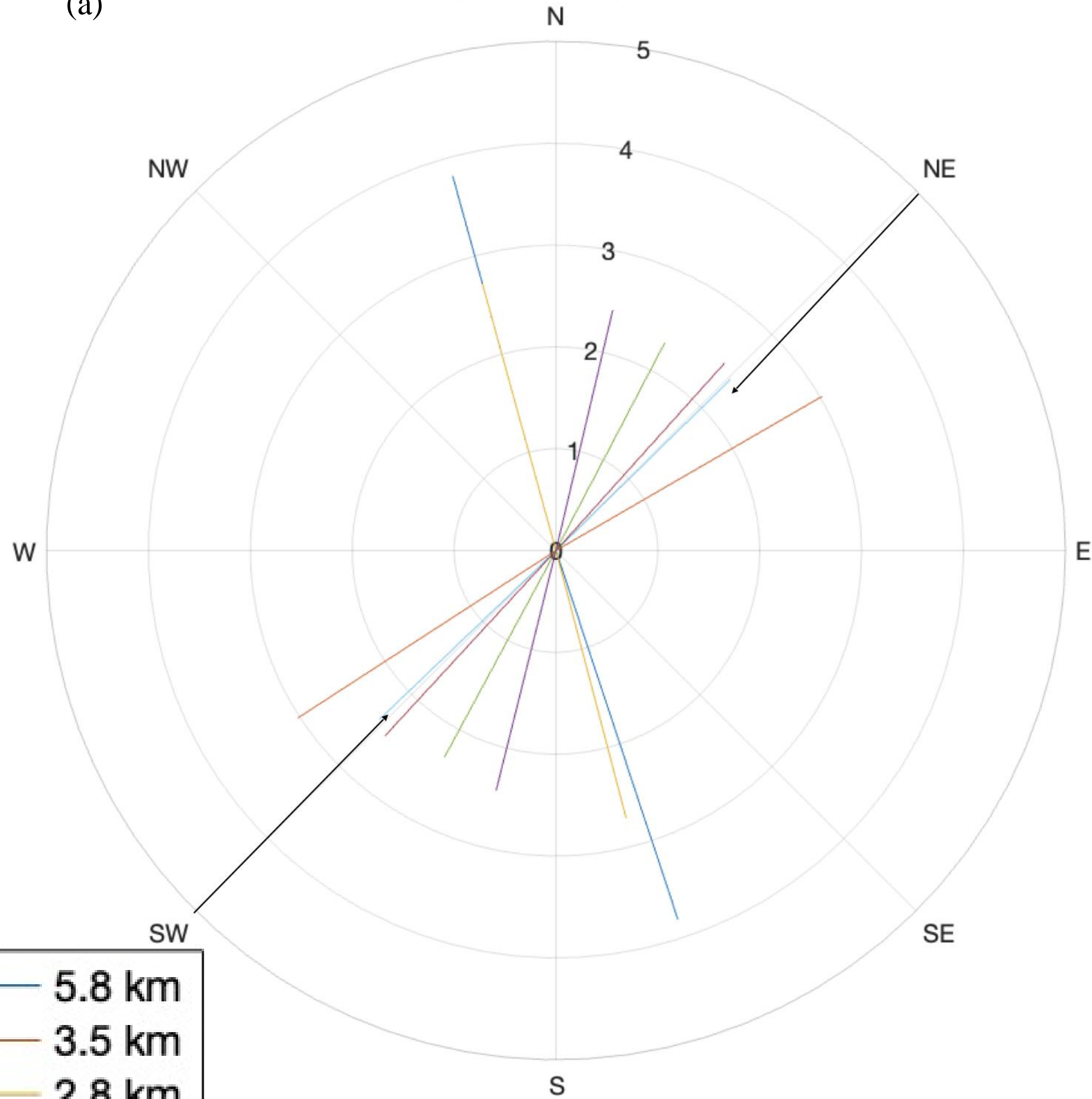
(b) **Love Wave Sensitivity Kernels**



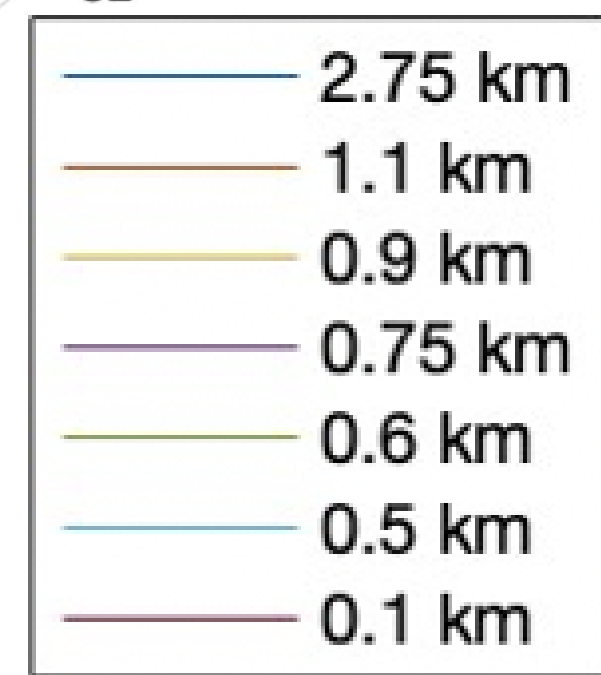
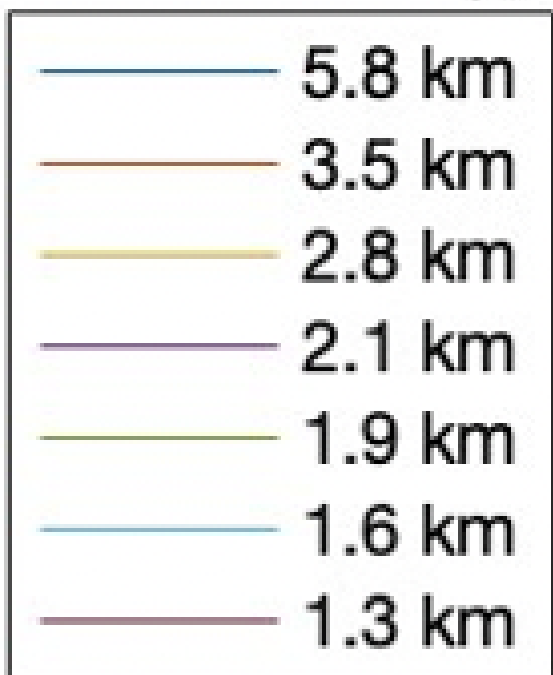
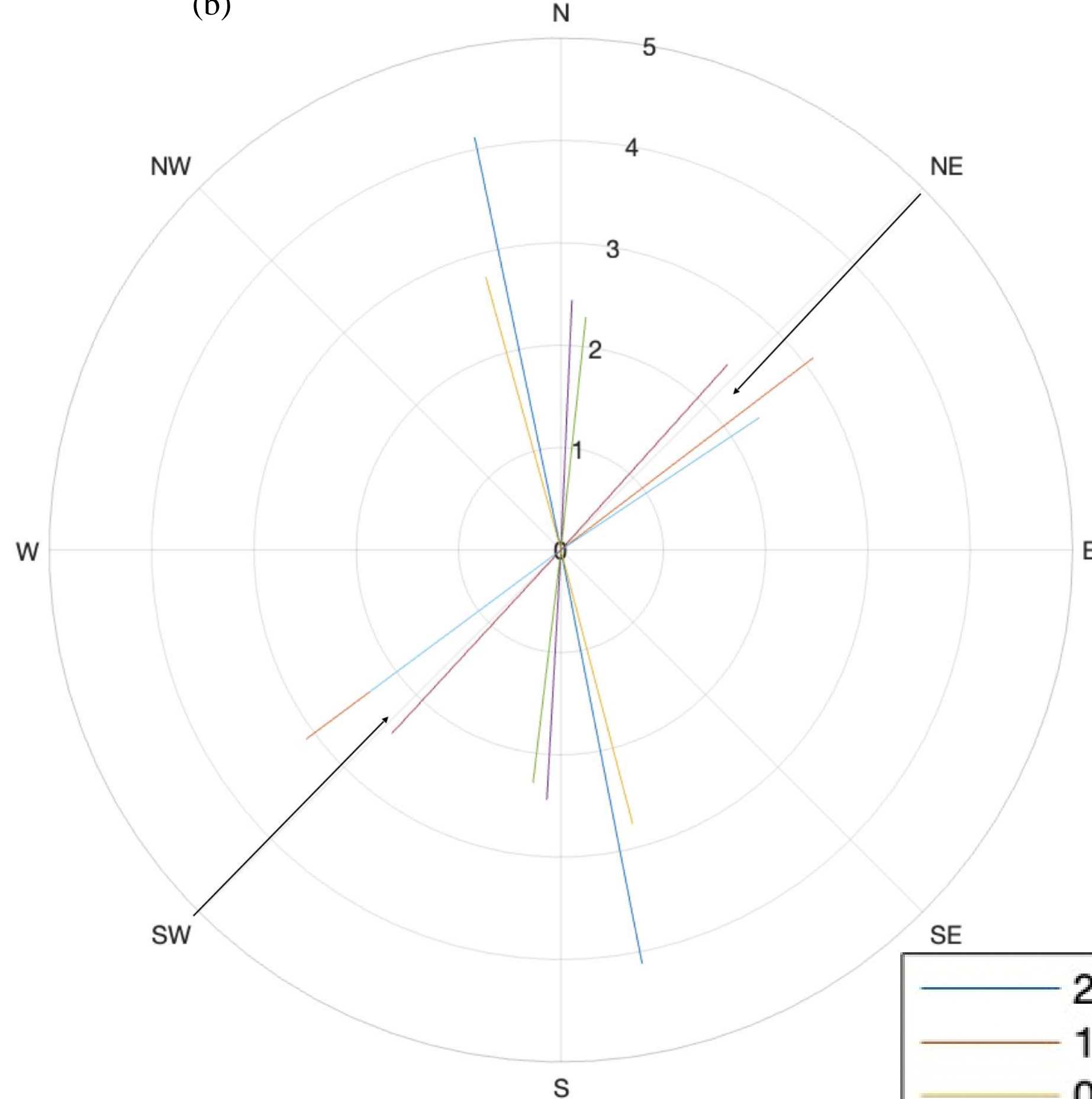
Depth Sensitivity

- 0.45 Hz
- 0.40 Hz
- 0.35 Hz
- 0.30 Hz
- 0.25 Hz
- 0.20 Hz
- 0.15 Hz
- 0.10 Hz

(a) Retrograde Rayleigh waves

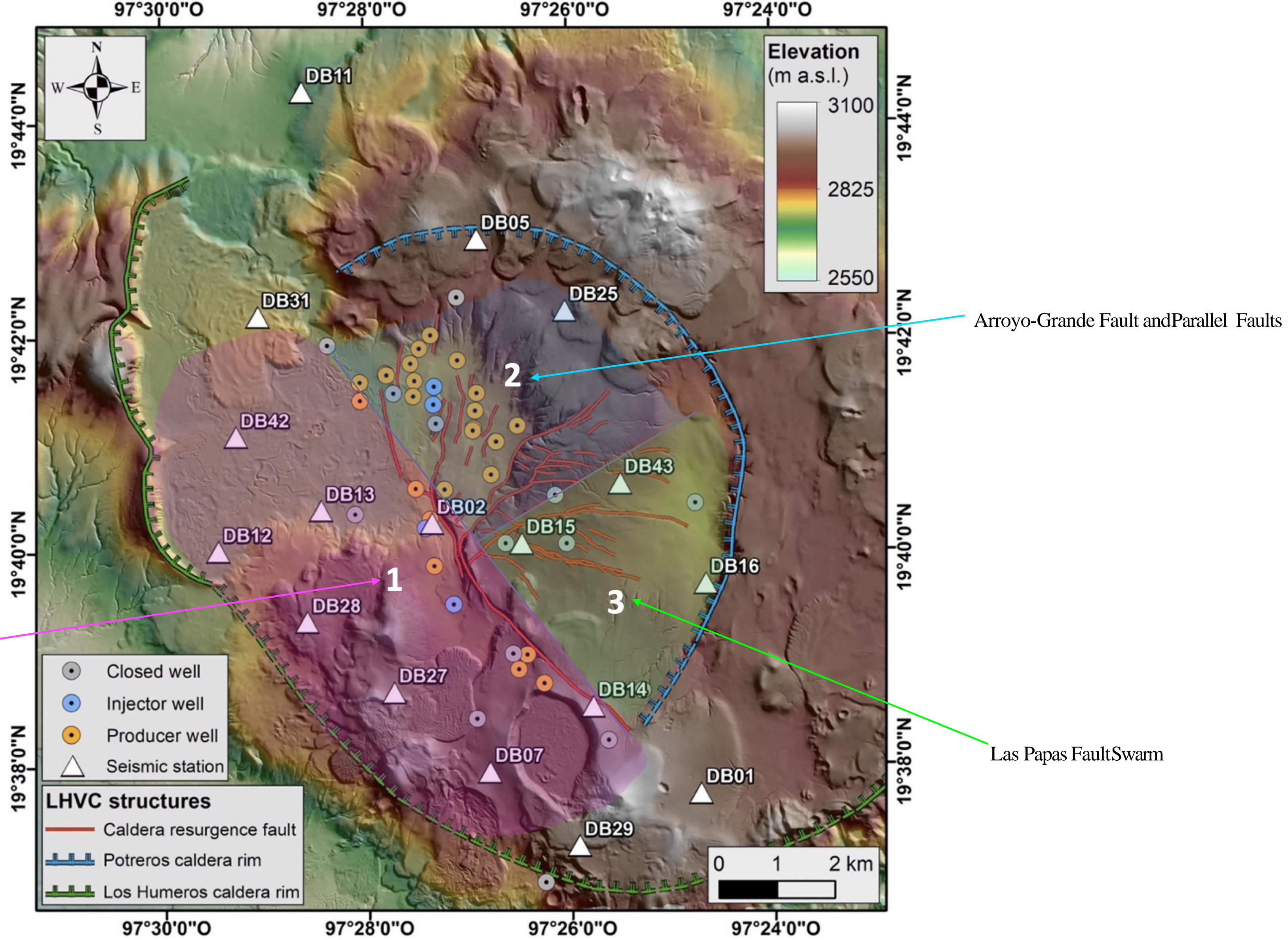


(b) Love waves





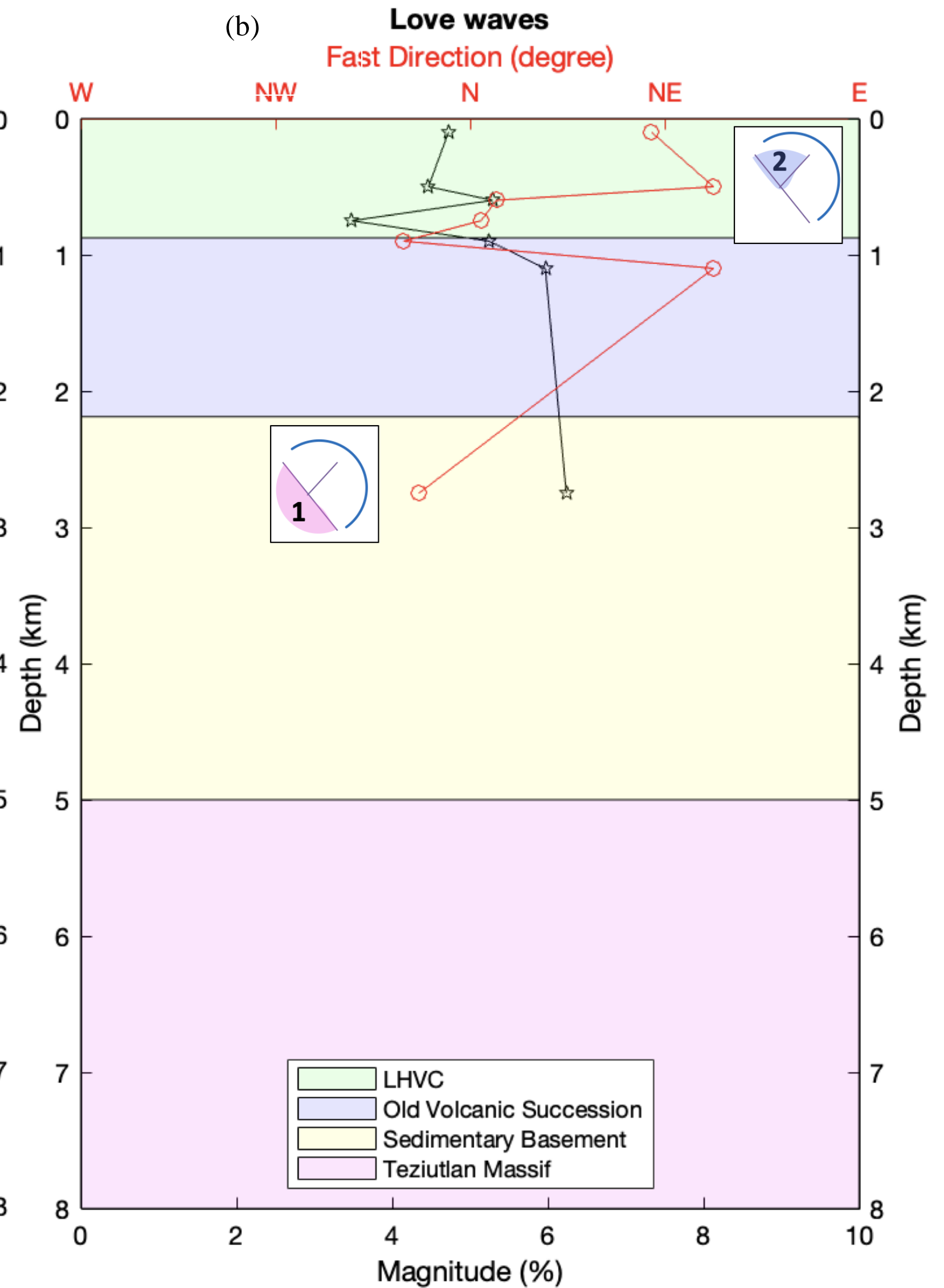
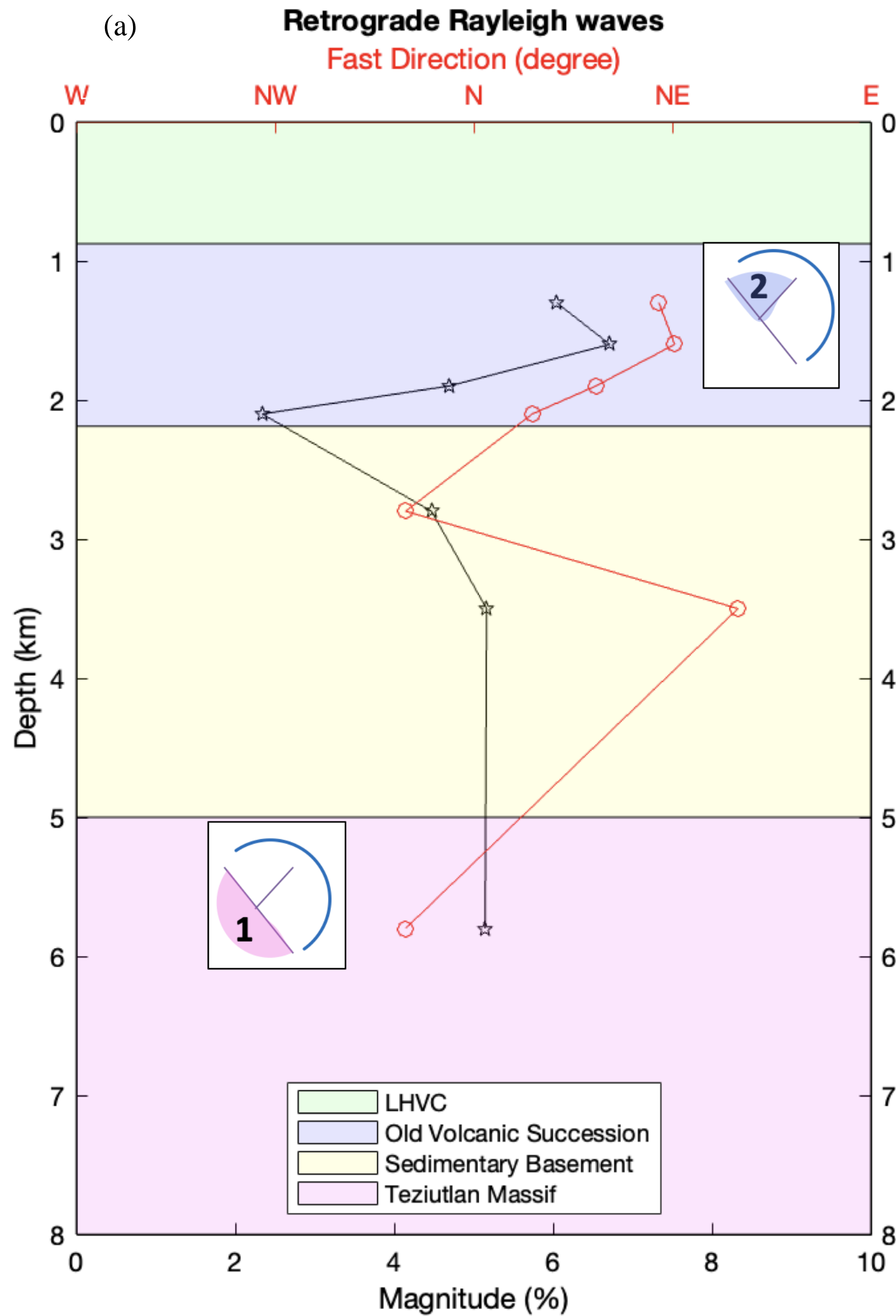
→ Regional Stress

Resurgence Faults and Fractures



Fast Direction varying with Depth

 Fast Direction
 Magnitude of Apparent Anisotropy

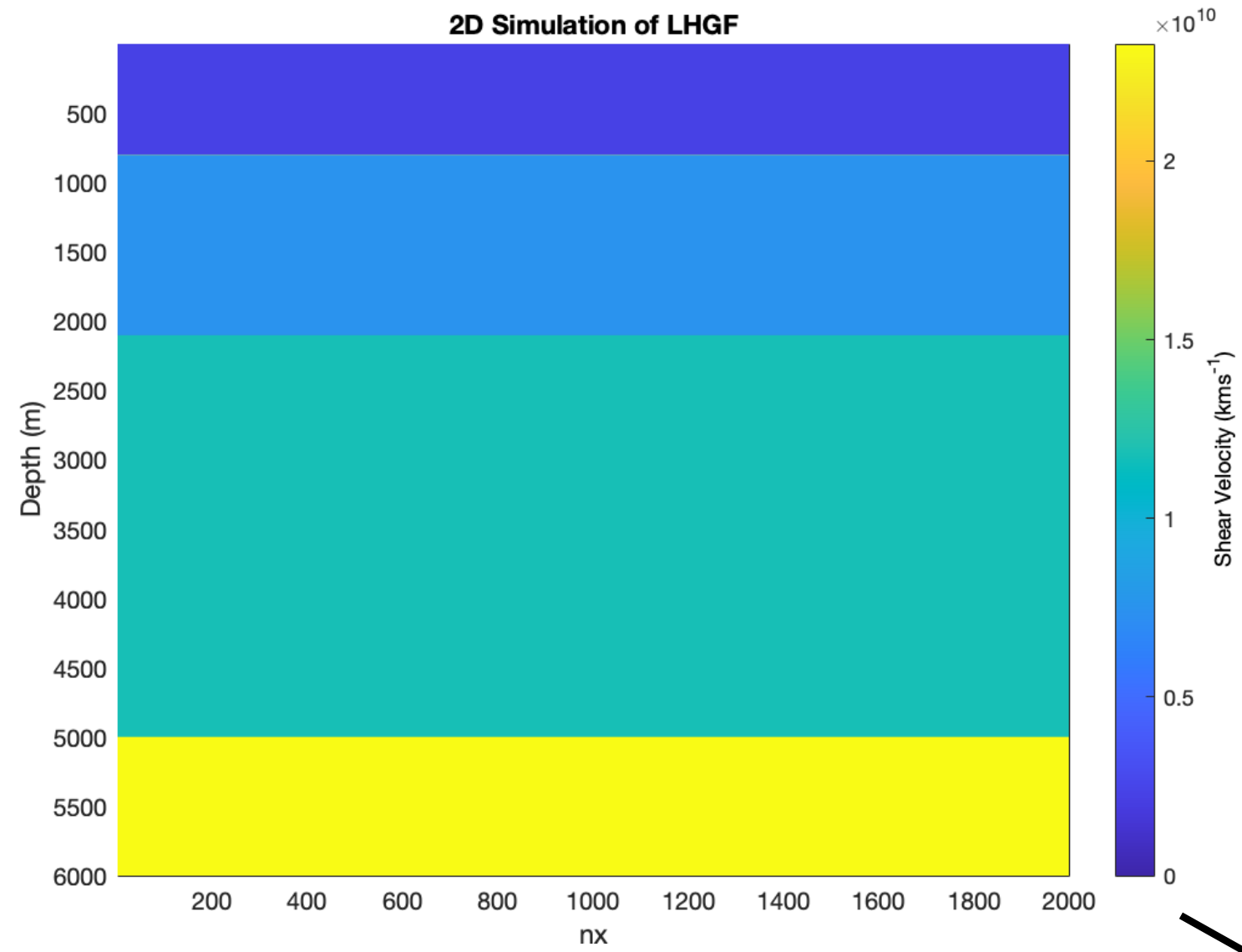


To Summarise

- 3C beamforming can estimate seismic anisotropy.
 - Clear fast directions from seismic anisotropy indicate subsurface structures.
 - Zone 1 and 2 dominant anisotropy match fast directions.
 - Maxtaloya-Los Humeros fault swarm (zone 1) and Arroyo Grande fault/parallel faults (zone) are, therefore, seen at depths > 2 km.
- NE-SW trending Mafic dykes may also have caused an anisotropy response (such as 3.5 km in
- Fig.6a).

Implications for the Geothermal Field:

- Continuation of faults at depth
- Anisotropy might be sensitive to hydrothermal productivity
- No evidence for brittle-ductile transition zone at depths < 6 km.



Future work - Simulation

Apply observed anisotropy

Wave Response

3C beamforming of simulated wave response

Compare to current results